Principles of structural geology and tectonic analysis

Strain

- In response to stress, rocks undergo transformations that are decomposed in 3 operations:
 - Rotation
 - Translation
 - Deformation which corresponds to a change in form
- Strain is a change in shape or internal configuration of a body

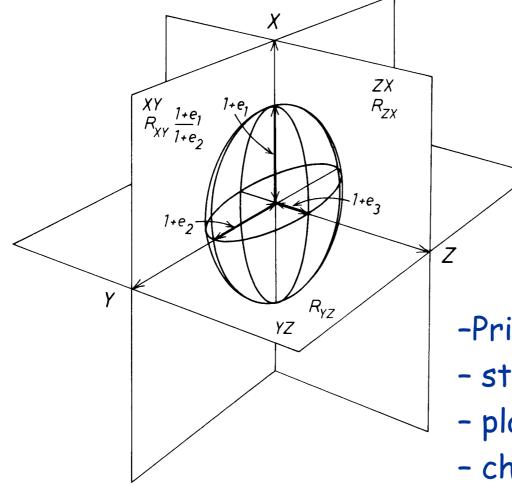
Strain

- Strain is marked by
 - Transformation of an initial circle/sphere into an ellipse/ellipsoid
 - The axis of deformation
 - Movement of particules defined by an instantaneous velocity vector, defining flow lines and finally trajectories
 - The kinematic notion, a function of time

Stress and strain

- Rheology is the study of the relationships between stress imposed on a body and resulting strains or strain rates
- Competence is a general term to describe the ease with which a material can deform
- In response to a stress, strain may be
 - Reversible, elastic deformation
 - Brittle, failure of a stressed body, at a certain state, the solid rocks lose their internal cohesion
 - Permanent, ductile flow without brittle deformation

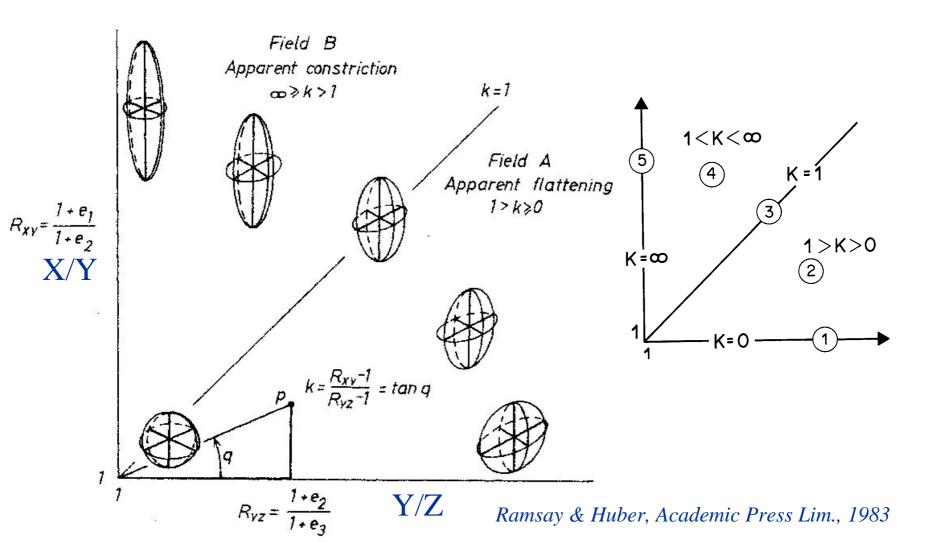
3 axis of deformation, a finite strain ellipsoid



-Principal

- strain directions, X, Y, Z
- planes XY, YZ, ZX
- changes in length e1, e2, e3
- ratios Rxy, Ryz, Rzx

The Flinn graph for representing the shape of any ellipsoid



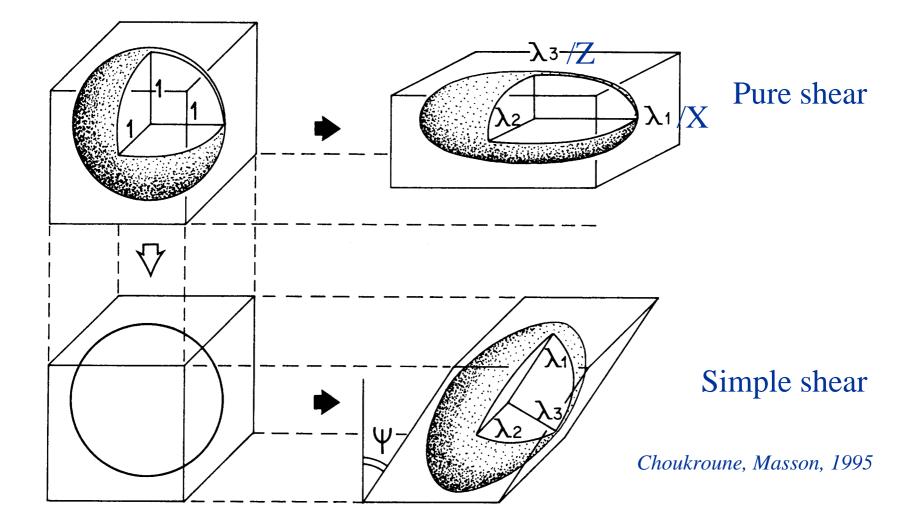
Strain regimes

- Deformation in rocks occurs under 2 strain regimes:
 - Pure shear
 - Simple shear
- Under combination of pure and simple shear

Simple shear

• A displacement which transforms an initial square in a parallelogram where the displacement vectors are all oriented parallel to one set of the opposite sides of the square and parallelogram. The vector direction is know as the shear direction.

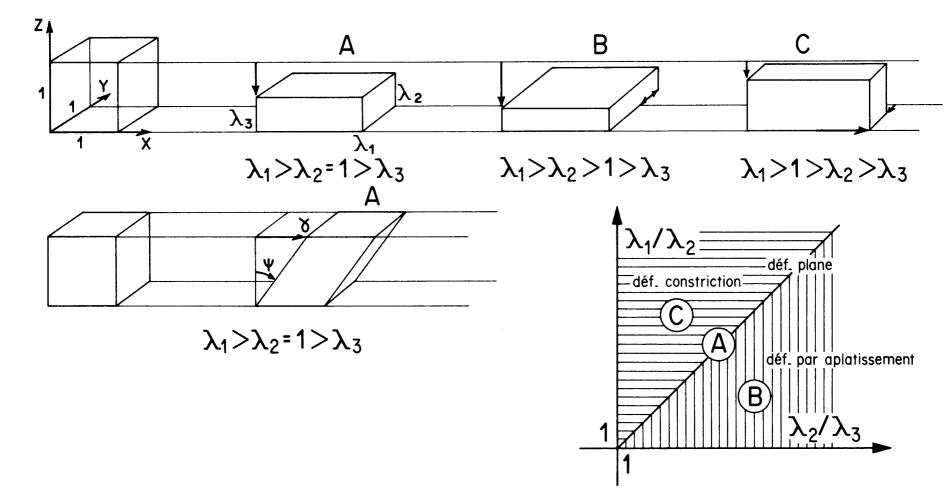
Transformation of a sphere in an ellispoid, 2 strain regimes



2 end-members

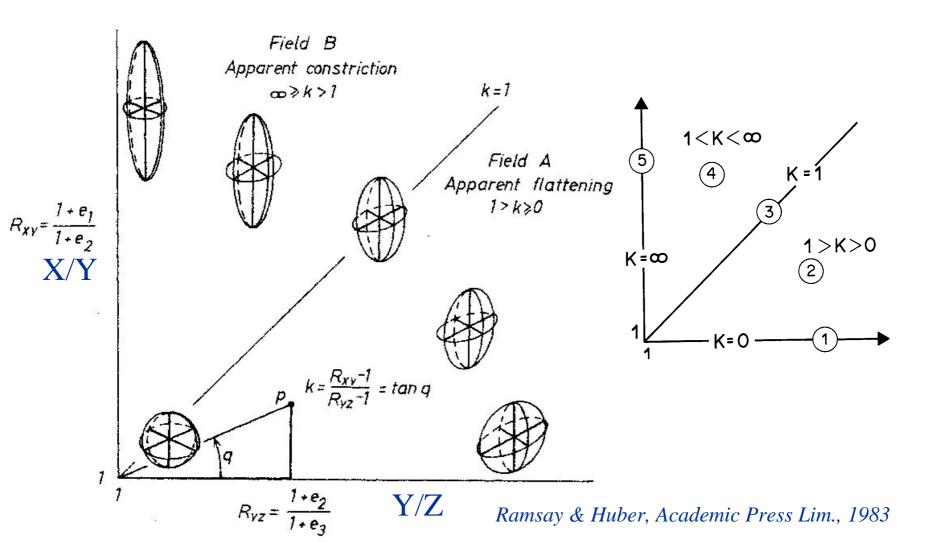
Pure shear	Simple shear
Axis of deformation remains parallel (coaxial	Axis of deformation rotate (non coaxial and
and non rotational)	rotational)
Flow lines tends to	Flow lines are parallel to
become perpendicular to	the shear direction during
the Z shortening axis	the deformation
Deformation >>	Deformation <<
Displacement	Displacement

Finite strain ellipsoid and strain regimes

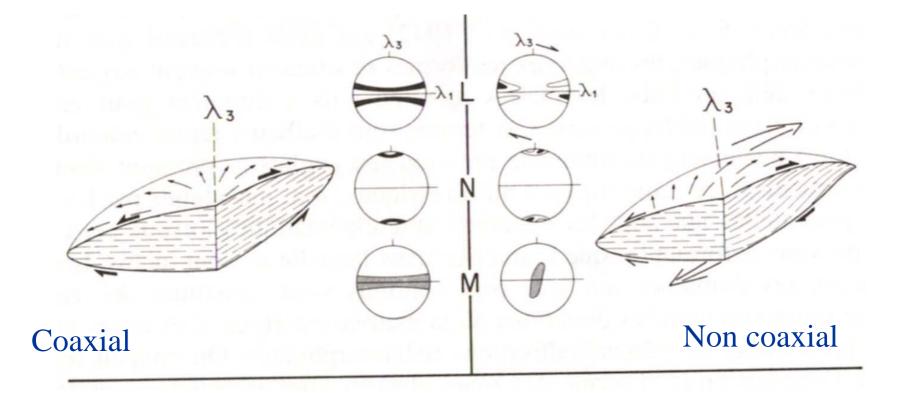


Choukroune, Masson, 1995

The Flinn graph for representing the shape of any ellipsoid



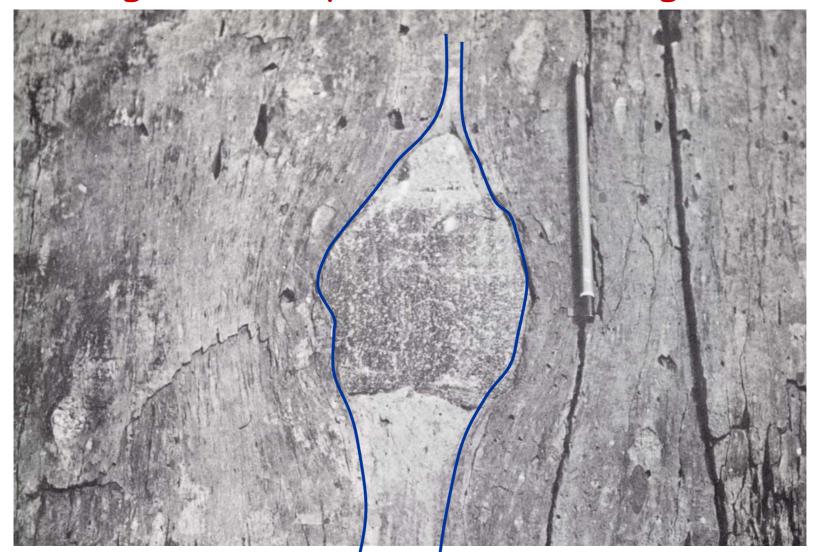
Change in geometry from coaxial to non coaxial regimes



Markers of the finite strain ellipsoid

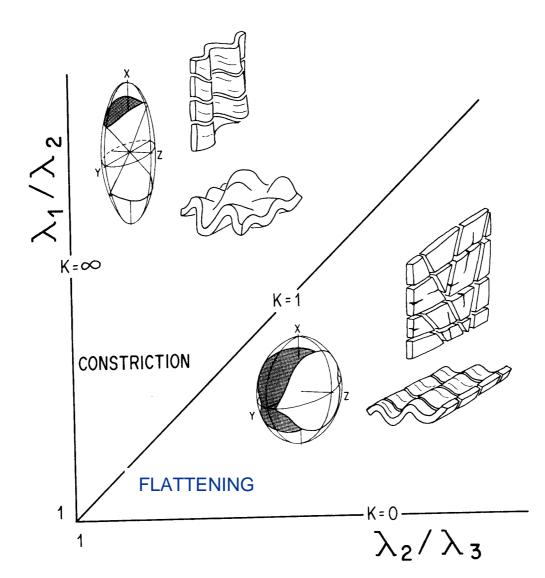
- Fold and boudinage structures
 - Boudinage is a structure produced during he extension of competent layers enclosed in an incompetent matrix

Deflection of cleavage around a competent fragment of quartzite in an argilite



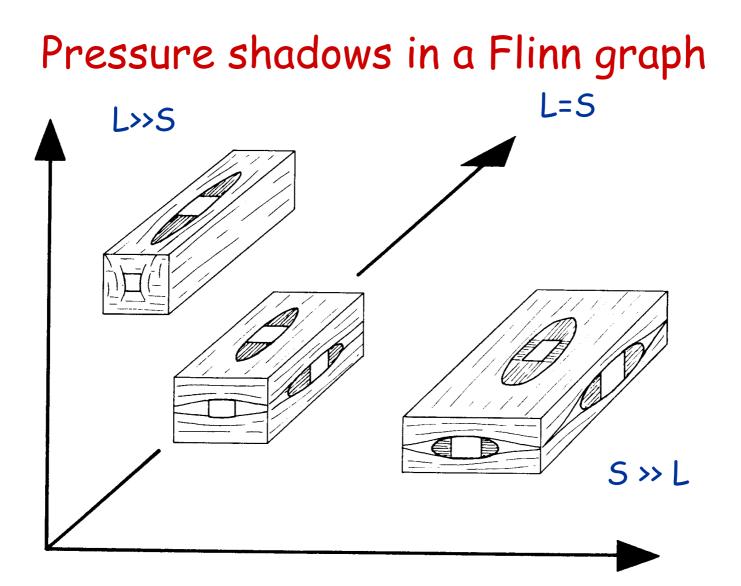
Ramsay & Huber, Academic Press Lim., 1983

Folding and boudinage in a Flinn graph



Markers of the finite strain ellipsoid

- Deformation of pre-tectonic objects
 - pebble, phenocryst, oolithe, redox spot, fossil..
- Mineral fabric and pressure shadows
 - Fabric is the geometric and spatial relationships between the crystal components making a rock, i.e. preferred orientation of grain shapes and crystallographic axis, the grain size...
 - Pressure shadow is a region of low strain protected from deformation by a rigid or competent object in a rock of lower competence



A linear type fabric: L>>S from the Hakos mountains

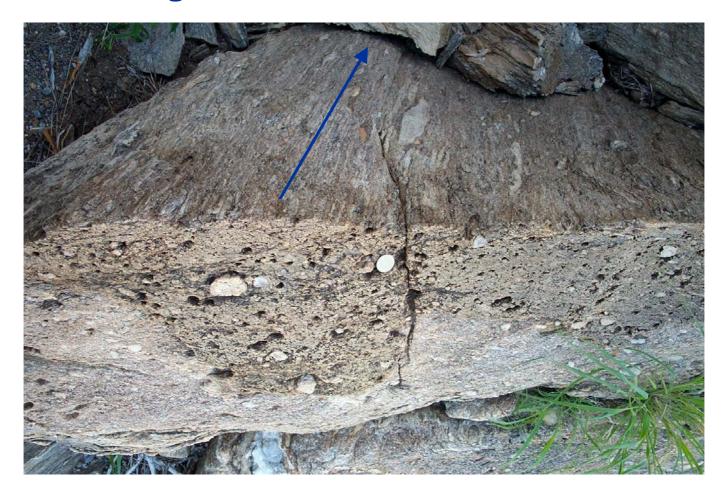
A granitic orthogneiss





A linear type fabric: L>>S from the Hakos mountains

A metaconglomerate



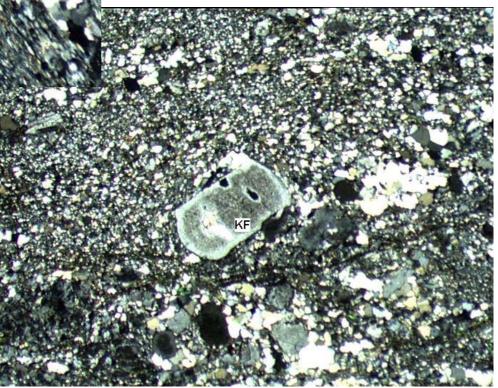


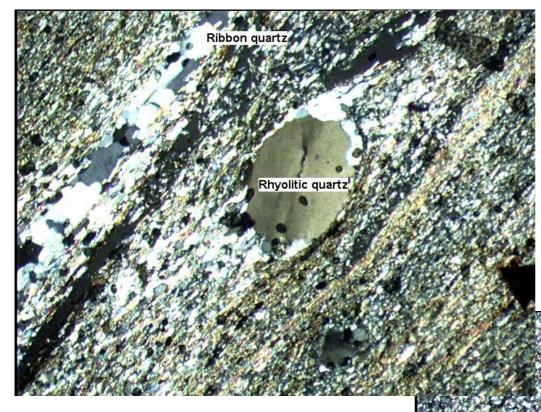
Sheared Gamsberg granite

XZ plane

1 single rock!

YZ plane





Sheared rhyolite

Rhyolitic quartz

XZ plane

1 single rock !

YZ plane

Partitioning and time...

Homogeneous vs non homogeneous deformation

- The definition of zones of homogeneous deformation
- The notion of bulk strain, an average strain in a heterogeneously deformed material
- Finite strain vs incremental deformation, notion of progressive deformation
 - Increment from one stage to the next
 - Rate of change along the main strain axis and of rotation defines the principal strain rates and the vorticity
 - Definition of the strain regime by identification and analysis of the markers of incremental deformation

Pure shear and development of structural elements

- Crenulation and mechanical rotation
- Pressure-solution, a deformation process whereby material under stress goes into solution at a localized point in a material, this material is transported by flow or diffusion
- Recrystallisation
- Transposition

Pure shear and development of folds , axial plane cleavage...

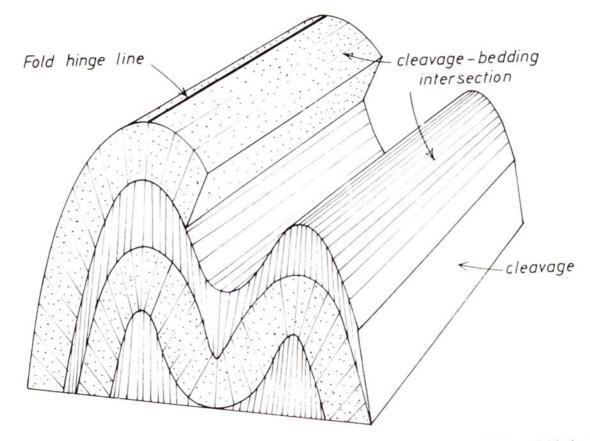


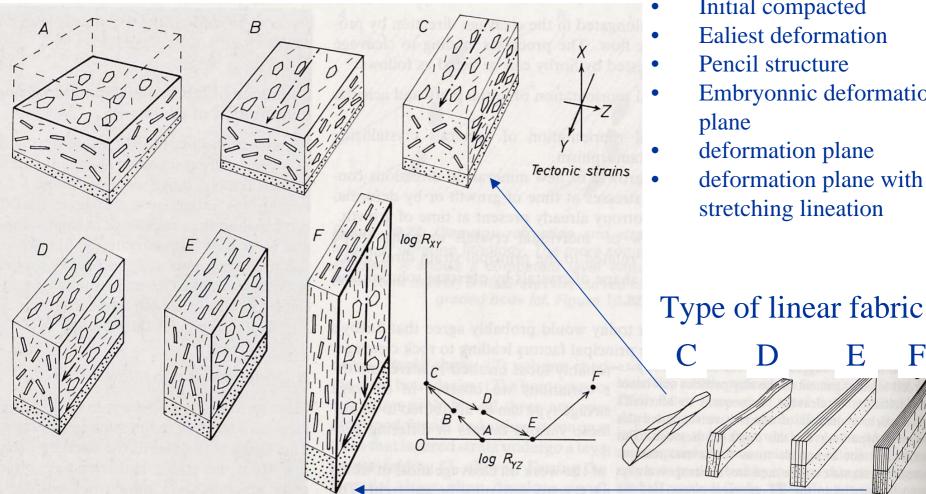
Figure 10.21. Geometric form of cleavage fans and cleavage–bedding intersection lineations in folded rocks. The most competent layers are stippled, the least competent layers unstippled. In simply folded rocks the cleavage–bedding intersection lineation is parallel to the fold hinge lines.

Pure shear and development of folds, axial plane cleavage...



Figure 10.18. Cleavage fans in folded Ordovician sediments, Rhosneigr, North Wales.

Progressive stages in fabric development of a shale



- Initial compacted
- Ealiest deformation
- Pencil structure
- Embryonnic deformation plane
- deformation plane
- deformation plane with stretching lineation

F

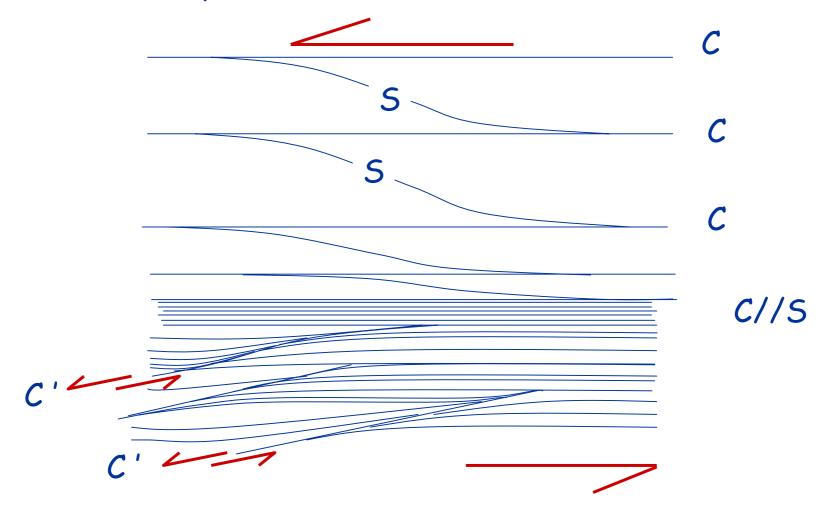
E

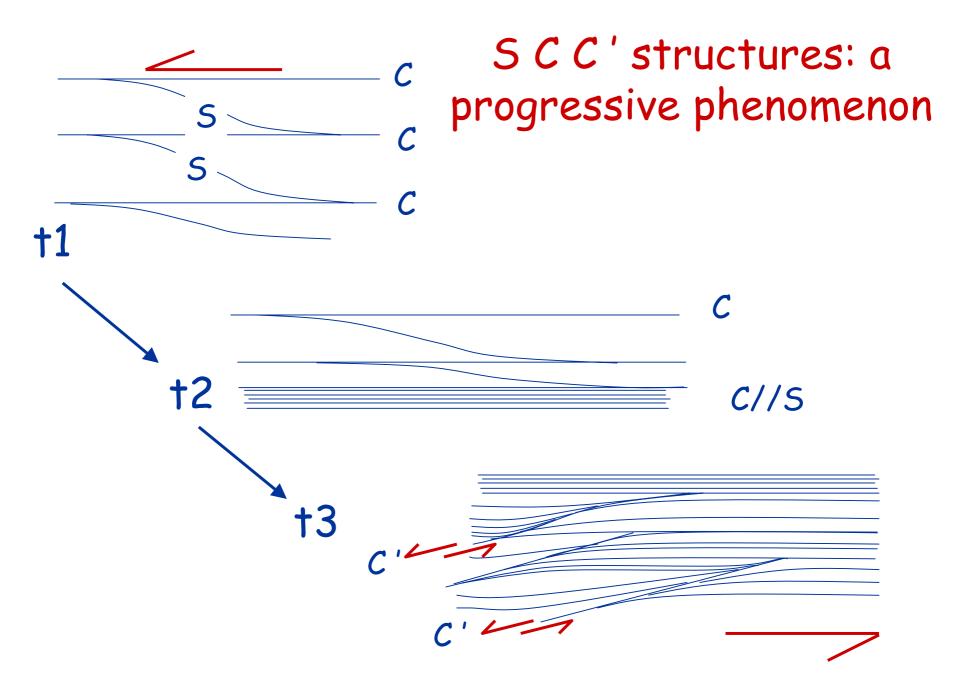
Simple shear and development of shear zone

- A shear zone is a zone with sub-parallel walls in which high deformations are localized
- Effects of non coaxial deformation
- Plane strain ellipsoid
- Markers of rotational deformation
 - Coexistence of deformation and displacement planes, S/C/C' structures
 - Assymetric pressure shadows and mineral fabric
 - Folding

S C C' structures

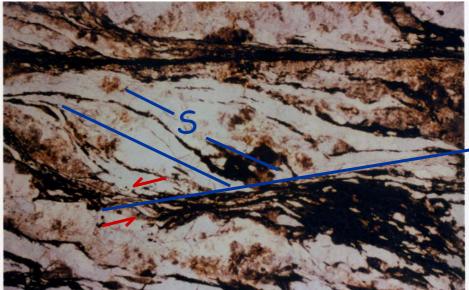
The S plane: the deformation plane, with no displacement
The C and C' planes: the displacement planes with high deformation (mylonitisation)



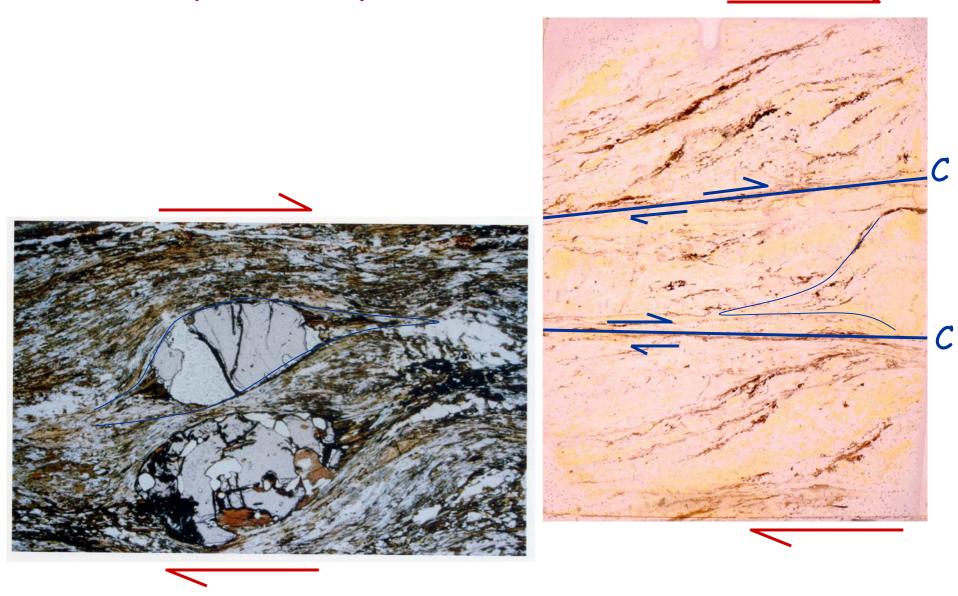


S C C' structures





Assymetric pressure shadow



Assymetric pressure shadow



Cleavage bending and helicitic inclusions

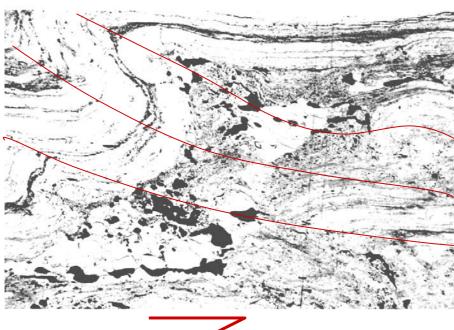




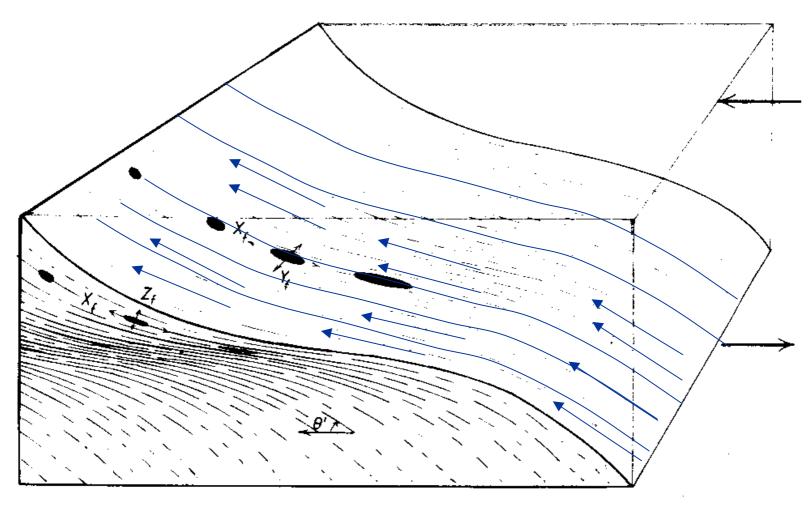


Assymetric folds and pressure shadows

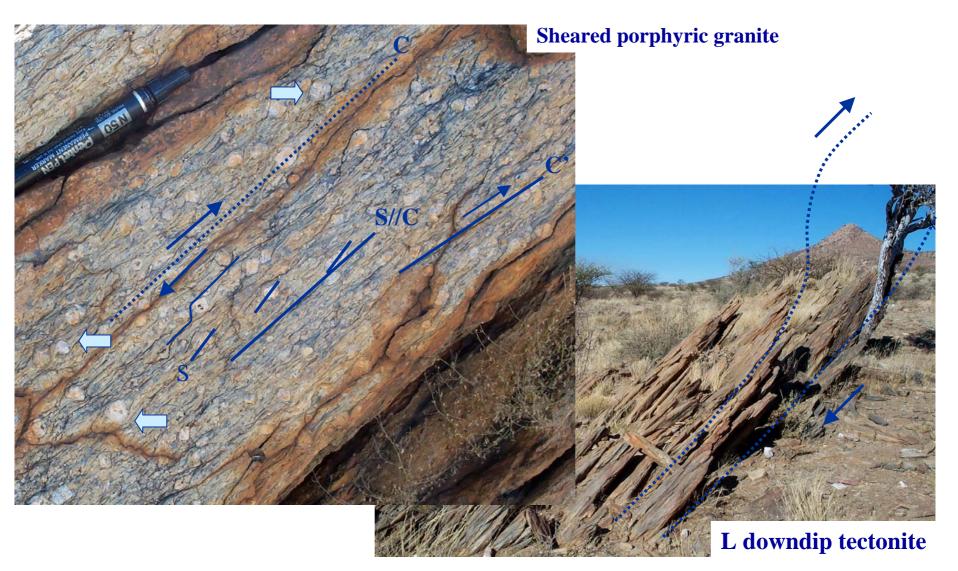




Fabric in a ductile simple shear zone



A reverse shear zone within the basement inlier of the Damara foreland



Development of non cylindrical folds during non coaxial deformation

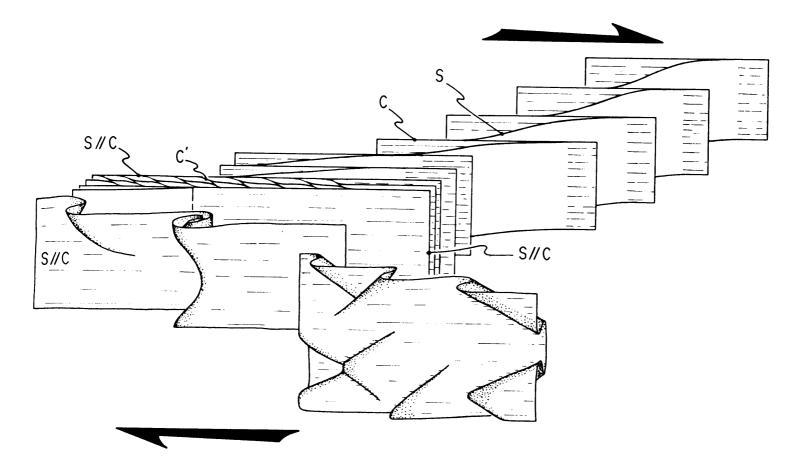


Fig. 147 – Schéma montrant l'évolution de la déformation dans la bordure sud du massif granitique impliqué dans le ciaillement sud-armoricain. C'est une séquence de mylonitisation avec les structures C/S.

Development of non cylindrical folds during non coaxial deformation

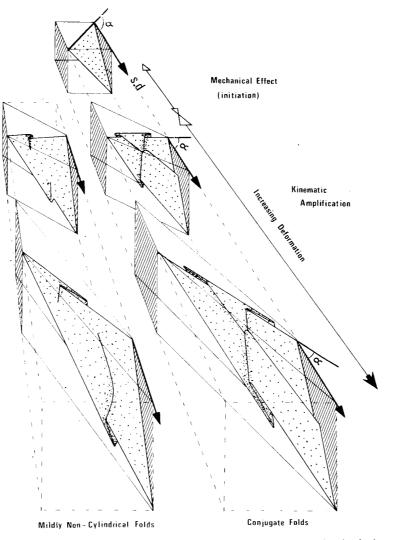
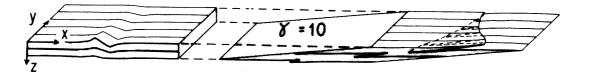


Fig. 6. Purely geometrical model for the development of the two types of fold during progressive simple shear (plane with diagonal shading is the shear plane). If the surface to be folded makes an initial angle (α) of more than 90° with the shear direction (s.d), conjugate fold sets develop (right). If the initial angle (α) is less than 90°, only one set develops (left).

Development of non cylindrical folds during non coaxial deformation



Progressive developement of sheath folds



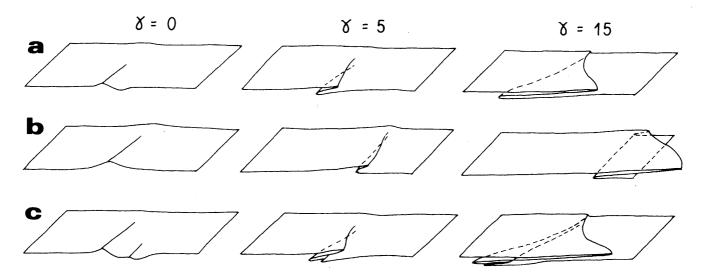
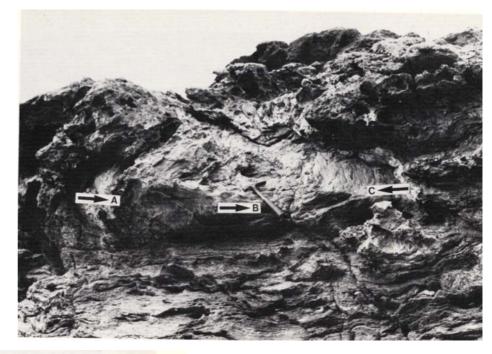


Fig. 2. Model 1: theoretical results. Noncylindrical initial deflections of various shapes (a, b & c) have been deformed in simple shear (top: x = shear direction, xy = shearing plane). Folds become asymmetric at moderate shear strains ($\gamma = 5$) and strongly noncylindrical at high shear strains ($\gamma = 15$).

Sheath folds from the Ile de Groix





Sheath folds from the base of a thrust nappe from the Variscan belt of France

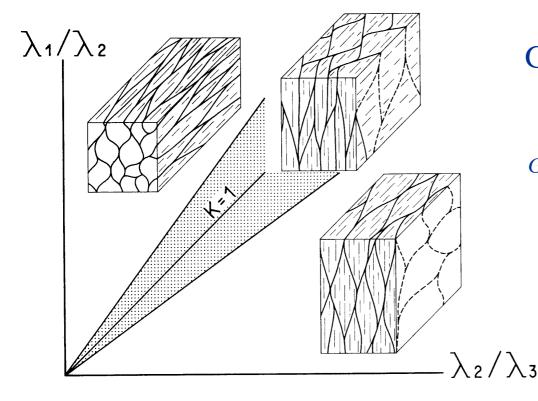




Sheath folds from the base of a thrust nappe from the Hakos mountains

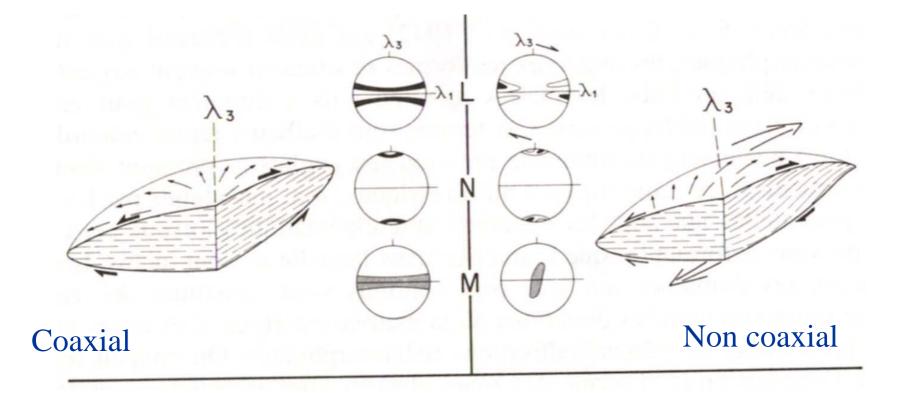


Shear zones and shear zones...

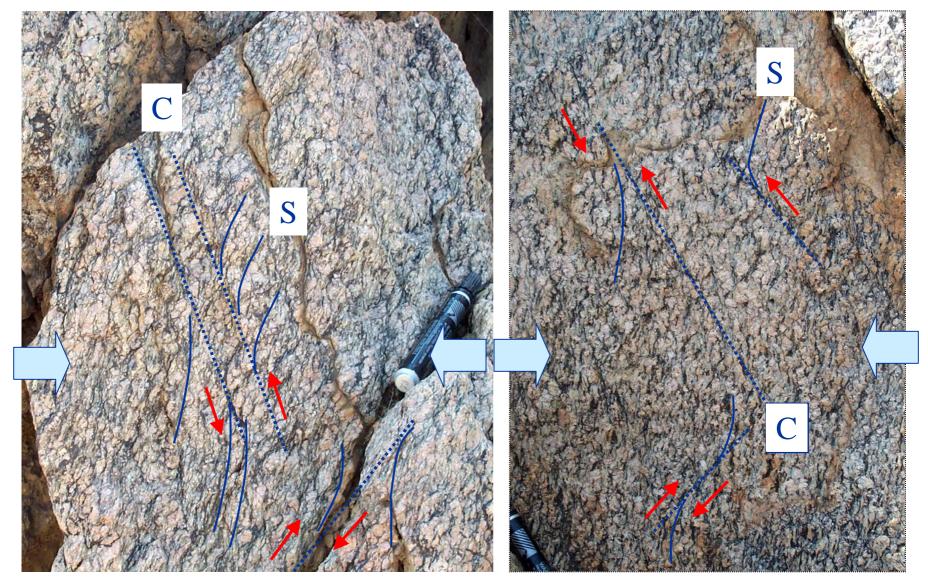


Geometry of the shear zones developed in flattening or constriction field *Gapais et al., Journ. Struct. Geol, 7, 1987*

Change in geometry from coaxial to non coaxial regimes



Conjugate shear zones: a coaxial regime SSE



vs a non coaxial regime...

